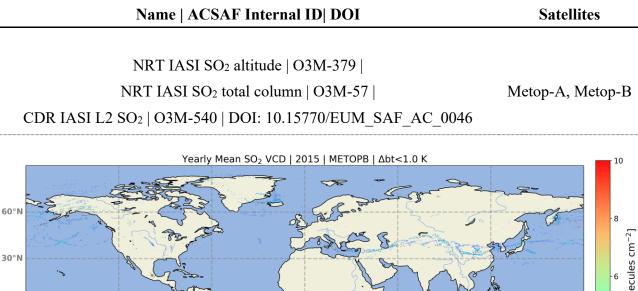


## ACSAF validation report on the EUMETSAT IASI/Metop CDR SO<sub>2</sub> datasets

Validated products:



60°N 30°N 40°N 

Yearly mean SO<sub>2</sub> vertical column density for 2015 including the Calbuco, Chile, eruption and a Mt Etna, Italy, eruption which transported SO<sub>2</sub> over Asia.

Authors: Name	Institute
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Gaia Pinardi	Belgian Institute for Space Aeronomy
Reporting period:	IASI/MetopA   Oct 2007 – Oct 2021
	IASI/MetopB   Feb 2013 - Dec 2021
Input data versions:	IASI/MetopA & /MetopB EUMETSAT netcdf



Data processor versions:

Brescia v6.6

### Introduction to EUMETSAT Satellite Application Facility on Atmospheric Composition monitoring (AC SAF)

#### Background

The monitoring of atmospheric chemistry is essential due to several human caused changes in the atmosphere, like global warming, loss of stratospheric ozone, increasing UV radiation, and pollution. Furthermore, the monitoring is used to react to the threats caused by natural hazards as well as follow the effects of the international protocols.

Therefore, monitoring the chemical composition and radiation of the atmosphere is a very important duty for EUMETSAT and the target is to provide information for policy makers, scientists and the general public.

#### Objectives

The main objectives of the AC SAF is to process, archive, validate and disseminate atmospheric composition products ( $O_3$ ,  $NO_2$ ,  $SO_2$ , BrO, HCHO,  $H_2O$ , OCIO, CO, NH3), aerosol products and surface ultraviolet radiation products utilising the satellites of EUMETSAT. The majority of the AC SAF products are based on data from the GOME-2 and IASI instruments onboard Metop satellites.

Another important task besides the near real-time (NRT) and offline data dissemination is the provision of long-term, high-quality atmospheric composition products resulting from reprocessing activities.

#### Product categories, timeliness and dissemination

*NRT products* are available in less than three hours after measurement. These products are disseminated via EUMETCast, WMO GTS or internet.

- Near real-time trace gas columns (total and tropospheric O<sub>3</sub> and NO<sub>2</sub>, total SO<sub>2</sub>, total HCHO, CO) and high-resolution ozone profiles
- Near real-time absorbing aerosol height and absorbing aerosol index from polarization measurement detectors
- Near real-time UV indexes, clear-sky and cloud-corrected

*Offline products* are available within two weeks after measurement and disseminated via dedicated AC SAF web services.

- Offline trace gas columns (total and tropospheric O<sub>3</sub> and NO<sub>2</sub>, total SO<sub>2</sub>, total BrO, total HCHO, total H<sub>2</sub>O) and high-resolution ozone profiles
- Offline absorbing aerosol height and absorbing aerosol index from polarization measurement detectors
- Offline surface UV, daily doses and daily maximum values with several weighting functions

Data records are available after reprocessing activities from the AC SAF archives.

- Data records generated in reprocessing
- Lambertian-equivalent reflectivity
- Total OClO

Users can access the AC SAF offline products and data records (free of charge) by registering at the AC SAF web site.

#### More information about the AC SAF project, products and services: <u>https://acsaf.org/</u>

AC SAF Helpdesk: <u>helpdesk@acsaf.org</u> Twitter: <u>https://twitter.com/Atmospheric\_SAF</u>



## **Applicable documents**

[ATBD] ALGORITHM THEORETICAL BASIS DOCUMENT, IASI Brescia SO2 columns and SO<sub>2</sub> altitude, Clarisse, L., Daniel Hurtmans, Pierre-François Coheur, Rosa Astoreca, Maya George and Cathy Clerbaux, SAF/AC/ULB/BresciaSO2\_ATBD, issue: 1.3, date: 01/02/2021.

[PRD] Product Requirements Document SAF/AC/FMI/RQ/PRD/001 Issue 2.0, 02/12/2022, https://acsaf.org/team/doc/prd/saf ac prd i2r0 20221202.pdf, last accessed: 17.03.2023.

[PUM] Product User manual, IASI Reprocessed SO<sub>2</sub> CDR (O3M-540), R. Astoreca, D. Hurtmans, L. Clarisse, P. Coheur, J. Hadji-Lazaro, M. George, C. Viatte, C. Clerbaux, M. Doutriaux-Boucher, C. Vicente and M. Crapeau, SAF/AC/ULB/PUM/010, 1.2, 14/09/2023.

[VR] IASI-C ACSAF VALIDATION REPORT, Koukouli, M. E., Pinardi, G. and Langerock, B., SAF/AC/AUTH/VR/IASI\_SO2 01/2021, issue: 01/2021, date: 30/4/2021, https://acsaf.org/docs/vr/Validation\_Report\_IASI-C\_SO2\_Apr\_2021.pdf, last accessed: 03.04.2023

## Authors

Gaia Pinardi, BIRA, authored section 3.2, while MariLiza Koukouli, AUTH, is responsible for the rest of the document. The summary and recommendations section is approved by both authors.

## Acknowledgments

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## **1. INTRODUCTION**

The main aim of this validation report is to testify as to the capability of the EUMETSAT operational algorithm v6.6 to accurately retrieve IASI/MetopA and /MetopB total SO<sub>2</sub> columns and SO<sub>2</sub> plume heights on a global scale.

The validation exercise is performed on orbital netcdf files created by EUMETSAT for the needs of the ACSAF team, communicated via a sftp location. Inherent in the netcdf file metadata is the following pertinent information:

summary = "This file contains the Release 1 IASI L2 reprocessed for one orbit. The reprocessing was done on Linux using an adapted version of the EUMETSAT operational algorithm V6.6. Until December 2016, the reprocessed Metop-A IASI L1C input data were used (doi:10.15770/EUM\_SEC\_CLM\_0014), after this date and for Metop-B the operational IASI L1c were used; this is to produce an homogeneous CDR. ERA5 were used as auxiliary model data"

processing centre = "CDR"

 $processing_mode = "R"$ 

project = "This data is provided as part of the EUMETSAT Satellite Application Facility on Atmospheric Composition Monitoring service, AC-SAF"

id = "DOI: 10.15770/EUM SAF AC 0046"

*history* = "*Release 1*"

*title* = "IASI SO2 CDR Release 1"

*title short name = "IASI SO2 CDR"* 

keywords = "IASI SO2 CDR Climate"

The time span of the delivered datasets is:

IASI/MetopA | 01 Jan 2008 to 15 Oct 2021

IASI/MetopB | 20 Feb 2013 to 31 Dec 2021



## **2. IASI/METOP SULPHUR DIOXIDE COLUMN AND PLUME HEIGHT**

## 2.1 The IASI/Metop CDR SO<sub>2</sub> products

The Brescia retrieval scheme includes two algorithms. The first algorithm, for the sounding of  $SO_2$  above ~5 km altitude, includes as main features a wide applicable total column range (over 4 orders of magnitude, from 0.5 to 5000 D.U.), a low theoretical uncertainty (3-5 %) and near real time applicability. The output consists of 5 different hypothetical SO<sub>2</sub> columns, for this version at 7, 10, 13, 16 and 25 km, depending on the altitude of the SO<sub>2</sub> plume in question, and is meant to be used with auxiliary altitude data. A typical use-case is where a user will have an estimate of the SO<sub>2</sub> altitude (provided by e.g. a model or third party measurements), and in that case, the user can interpolate the 5 SO<sub>2</sub> columns to this estimate to obtain a single SO<sub>2</sub> column.

These different SO<sub>2</sub> columns which depend on different SO<sub>2</sub> plume heights have already been validated using previous versions of the IASI datasets, in the official validation reports for <u>IASI-A</u> and <u>IASI-B</u> and for <u>IASI-C</u>. In the current validation report, we will focus on the retrieval by the second algorithm where information on the SO<sub>2</sub> altitude can also be retrieved directly from the IASI measurement. As a result, an interpolated SO<sub>2</sub> column (from the first algorithm), at the retrieved SO<sub>2</sub> altitude (from the second algorithm) is also provided. This SO<sub>2</sub> column product, and the associated retrieved SO<sub>2</sub> altitude are validated in this report.

Refer to the [PUM] and the [ATBD] for further information on the IASI Brescia products.

#### 2.1.1 SO<sub>2</sub> vertical column density and SO<sub>2</sub> altitude

In this work, the  $so2\_col$  variable included in the netcdf files provided by EUMETSAT is being assessed, which represents the  $SO_2$  column at the retrieved plume altitude from a neural network approach, as well as the  $so2\_altitudes$  variable which represents the retrieved plume altitude. Two further variables are also utilized, namely the  $so2\_aflag$ , general quality flags and the  $so2\_bt\_difference$ , which represents an indicative brightness temperature difference; it is recommended to only keep the data for which this difference is greater than 0.4.

With respect to the *so2\_qflag*, it is stated in the [PUM] that all retrieved SO<sub>2</sub> columns are considered best quality retrievals and can be used. However, when the *so2\_qflag* is set 9 this identifies retrievals which used the IASI L2 pressure and temperature profiles. When the *so2\_qflag* is set to either 11, this means that the pressure and temperature profiles are missing in the IASI L2 data and that model/forecast data have been used instead. Pixels with associated *so2\_qflag* set to zero identify missing pixels. In the following, pixels retrieved with the IASI L2 temperature and pressure profiles were allowed for the analysis.

With respect to the *so2\_bt\_difference*, it is stated in the [PUM] only look at the retrievals in the neighbourhood of  $\Delta$ Bt>1K pixels, and not use at all any pixels with an associated  $\Delta$ Bt <0.4K, which means that there is not enough SO<sub>2</sub> to have a reliable retrieval. It is further stated that size of the region around  $\Delta$ Bt>1K pixels is left to the user's estimate. This recommendation allows for several possible mis-interpretation scenarios. We hence investigated using  $\Delta$ Bt>0.4K as flag and, based on the findings of Section 2.2.1, we have used only pixels with associated  $\Delta$ Bt>1.0K for the remainder of the report.



#### 2.1.2 Volcanic SO<sub>2</sub> sources

In order to assess the IASI/Metop CDR SO<sub>2</sub> VCDs, a number of volcanic eruptions were chosen, enumerated in Table 1. In the following document, a number of these eruptions will be used to demonstrate the various aspects of the dataset as well as in the validation against the correlative dataset, that of S5P/TROPOMI.

Starting day	Ending day	Name	Country
20071003	20071007	JebelalTair	Yemen
20080511	20080514	Etna	Italy
20080718	20080719	Okmok	Alaska
20080811	20080906	Kasatoschi	Alaska
20081105	20081112	Dalafilla	Ethiopia
20090619	20090716	Sarychev	Russia
20100515	20100520	Eyjafjallajoekull	Iceland
20110523	20110604	Grimsvotn	Iceland
20110618	20110627	Nabro	Eritrea
20111105	20111117	Niryacongo	Congo
20121223	20121231	Copahue	Chile
20140214	20140222	Kelut	Indonesia
20150422	20150430	Calbuco	Chile
20160102	20160115	Soputan	Indonesia
20160326	20160330	Pavlov	Alaska
20180219	20180228	Sinabung	Indonesia
20180626	20180702	SierraNegra	Ecuador
20180726	20180802	Ambae	Vanuatu
20190622	20190715	Raikoke	Russia
20200112	20200115	Taal	Philippines
20211219	20211231	HungaTonga	Tonga

 Table 1. Selected volcanic emissions studied within this work.

### 2.2 Inter-sensor investigations

In the following section we first investigate the effect of the recommended filtering of the IASI/Metop CDR as well as the inter-sensor consistency of the two instruments.



2.2.1 Investigation on the  $\Delta Bt$  flag

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## 

**Figure 1.** The effect of the  $\Delta Bt$  recommended cut-off on the IASI SO<sub>2</sub> reported loads for three major eruptions during the IASI-A timeline. Left column,  $\Delta Bt > 0.4K$  and right column,  $\Delta Bt > 1K$ .

As discussed in Section 2.1.1, there exist a recommendation that pixels with associated  $\Delta Bt > 0.4K$  should be used if in the vicinity of pixels with associated  $\Delta Bt > 1K$ . The extend of this vicinity is left to the user to define. To begin with, we show in Figure 1, the effect of applying a  $\Delta Bt > 1K$  in the entire dataset [right column] on three major eruptions during the IASI-A timeline. All eruptive days are simply plotted overlain, without any gridding, sorted per SO<sub>2</sub> load so that the highest loads appear in the plot. In all cases, a number of high SO<sub>2</sub> loads are filtered out by the strict filter of  $\Delta Bt > 1K$ .

For both cases, however more pronounced for the Kasatochi eruption shown in the upper row of Figure 1, numerous high SO<sub>2</sub> loads are excluded when using the stricter filtering. We overall consider the recommendation of "choosing a vicinity" beyond the typical abilities of the standard user of satellite data and would suggest that a flag pointing to such occurrences be included in future versions of the dataset. The ACSAF GOME2/Metop SO<sub>2</sub> product (Valks et al., 2019) includes such a volcanic flag and a similar approach could be applied to the IASI/Metop SO<sub>2</sub> products.



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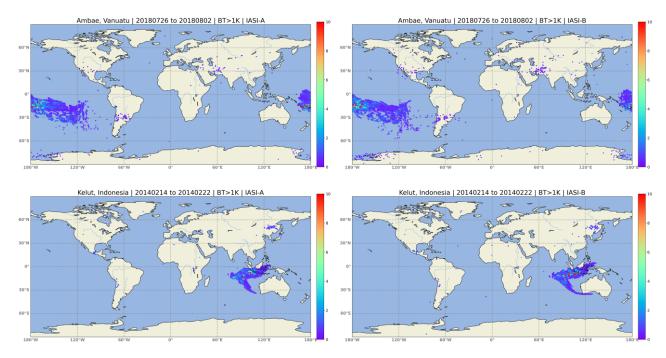
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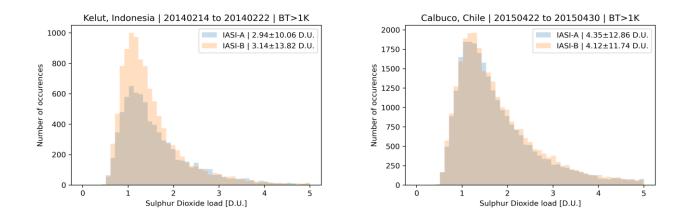
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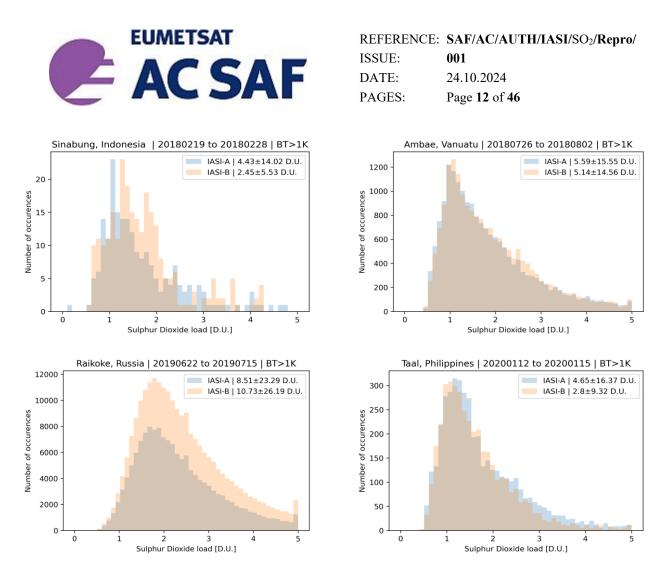
#### 2.2.2 Inter-sensor consistency

The inter-sensor consistency is examined using known volcanic eruptions during the tandem period of operations of the two instruments. In Figure 2, two such examples are presented for the Ambae, 2018 eruptive period [upper] and the Kelut, 2014, eruptive period [lower]. The locations and spreads of the plumes coincide to a significant degree, also demonstrated in the form of histogram representations of the SO<sub>2</sub> loads in Figure 3.



**Figure 2.** The IASI-A [left] and IASI-B [right]  $\Delta Bt > 1K SO_2$  loads during the Ambae, 2018 eruption [upper] and the Kelut, 2014 eruption [lower].





**Figure 3.** Histograms of the IASI-A [blue] and IASI-B [orange]  $\Delta Bt > 1K$  SO<sub>2</sub> loads during six concurrent volcanic eruptions between 2014 and 2020 in temporal sequence. Names and time periods are insert in the titles.

In Figure 3, histograms of the IASI-A [blue] and IASI-B [orange]  $\Delta Bt > 1K$  SO<sub>2</sub> loads during six concurrent volcanic eruptions between 2014 and 2020 are presented. The eruptions were chosen on purpose so as to identify possible issues that may affect users. The timelines of the eruptive days are shown in each title and the mean loads and 1-sigma deviation in the figure legend. The first issue to note is that differences appear to exist for the Kelut eruption [upper left], the Sinabung [middle left] and the Raikoke eruption [bottom left], which however are due to a few missing IASI-A days from the timeline. The second issue affects the Taal [bottom right] eruptions where both sensors reported data during all four eruptive days. However IASI-A reports a number of pixels with high SO<sub>2</sub> load which are either not sensed by IASI-B, or do not pass the recommended flag filtering or do not pass the  $\Delta Bt > 1K$  flag. This results in the difference in mean loads presented in this case. Finally, during eruptions were neither of these issues are present, such as during the Calbuco [upper right] and Ambae [middle right] eruptions, the mean loads and distributions between the two sensors are in agreement within the first decimal place.

We hence conclude that the IASI-A and IASI-B SO<sub>2</sub> CDR datasets can be merged and used in tandem, when the user prefers to increase the collocations, always considering the time difference of the two observations.

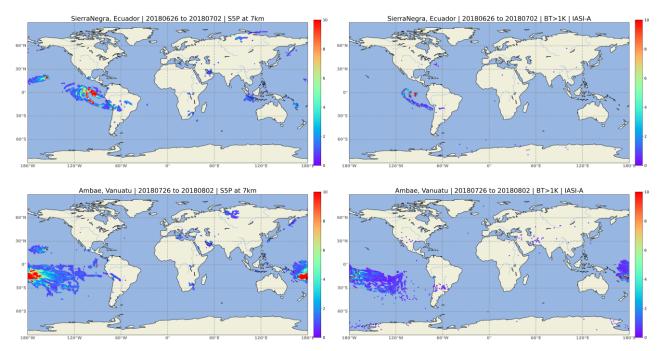


## **3.** SULPHUR DIOXIDE COLUMN VALIDATION

In this section, comparisons of the IASI/Metop CDR SO<sub>2</sub> columns are presented against collocated satellite estimates by S5P/TROPOMI and ground-based estimates by BIRA MAX-DOAS systems.

## 3.1 Validation against S5P/TROPOMI

The TROPOMI sensor on board the Sentinel-5P platform samples the Earth's surface with a revisit time of one day and a spatial resolution of  $5.5x3.5 \text{ km}^2$  since August 2019. This allows the resolution of fine details of SO<sub>2</sub> plumes and S5P is arguably a valuable tool in anthropogenic SO<sub>2</sub> emissions studies but also volcanic emissions, from degassing to eruptive processes. The SO<sub>2</sub> retrieval algorithm is based on the DOAS technique and is fully described in the relevant ATBD [Theys et al., 2022] as well as in Theys et al., 2017, where additional verification results as part of the algorithm development are also presented. The S5P/TROPOMI SO<sub>2</sub> dataset is publicly available via the Copernicus Open Access Hub, <u>Sentinel-5P Pre-Operations Data Hub (copernicus.eu</u>). In the following, offline data between 01.05.2018 and 31.12.2019, versions 01.01.05 and 01.01.06, were downloaded. As suggested in the ReadMe file<sup>1</sup>, in case of volcanic emissions the only filtering criteria required is that the solar zenith angle is < 70° and the sulfur dioxide detection flag to be raised. The S5P/TROPOMI volcanic SO<sub>2</sub> at an assumed 7km plume was gridded on a 0.25x0.25° grid on a daily basis. In Figure 4, the daily SO<sub>2</sub> load (D.U.) for three major eruptions, Sierra Negra and Ambae, in 2018, and Raikoke, in 2019 is shown on the left, and the equivalent IASI-A fields on the right column.



<sup>&</sup>lt;sup>1</sup> https://sentinels.copernicus.eu/documents/247904/3541451/Sentinel-5P-Sulphur-Dioxide-Readme.pdf

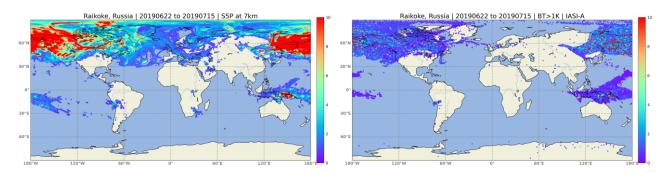


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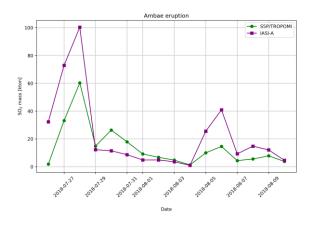


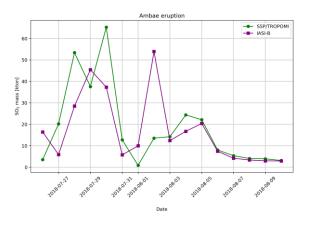
**Figure 4.** S5P daily volcanic SO<sub>2</sub> VCD at 7km for three major eruptions, Sierra Negra and Ambae, in 2018, and Raikoke, in 2019 shown on the left column and the IASI-A  $\Delta Bt$ > 1K observations for the same eruptions on the right column.

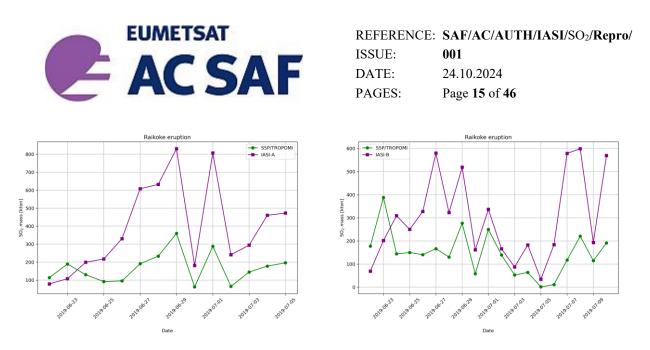
In Table 2, the mean volcanic SO<sub>2</sub> loads per eruption between S5P, IASI-A and IASI-B (where applicable) are provided for the three eruptions shown in Figure 4. The S5P and IASI observations are simply plotted overlain in these figures sorted in the manner that the maximum volcanic SO<sub>2</sub> loads appear on top. As a result, for the many days of each eruption, all of the higher loads [redder points] appear over the rest, hence giving the impression that S5P has the higher loads and higher variabilities. This is however not the case, as shown in the statistics provided in Table 2, which are based on the points plotted in in Figure 4. Overall, the mean IASI SO<sub>2</sub> loads report higher values, and a far higher variability – denoted by the 1-sigma of the mean – than the S5P SO<sub>2</sub> loads. The IASI-A and -B maxima are also far larger than the equivalent ones for S5P which led us to proceed in calculating the SO<sub>2</sub> masses in ktons/day, as a more representative comparative metric.

	S5P at 7km [D.U.]			IASI-A [D.U.]			IASI-B[D.U.]		
	min	max	mean±std	min	max	mean±std	min	max	mean±std
Sierra Negra	0.32	171.3	1.9±3.28	0.2	227.0	14.03±32.17	n/a	n/a	n/a
Ambae	0.34	195.0	2.16±4.11	0.2	250.0	5.11±16.52	0.2	314.0	4.22±14.13
Raikoke	0.29	802.0	2.98±8.26	0.2	1658.0	6.24±22.17	0.1	1581.0	6.8±22.6

Table 2. Comparison of the mean SO<sub>2</sub> load per eruption between S5P, IASI-A and IASI-B.







**Figure 5.** Same-pixel comparison of SO<sub>2</sub> mass [kt/day] for the Ambae [upper] and the Raikoke [lower] eruptions between S5P and IASI-A [left] and IASI-B [right].

In Figure 5, the collocated pixels between S5P and IASI-A [left] and IASI-B [right] were used to calculate the SO<sub>2</sub> masses in kt/day emitted after the Ambae 2018 (upper) and Raikoke 2019 (lower) eruptions. The two sensors report a nice co-variability in the days following the eruptions, with IASI reporting higher loads, as expected when considering the SO<sub>2</sub> loads shown in Table 2.

#### 3.2 Comparisons against the MAX-DOAS network

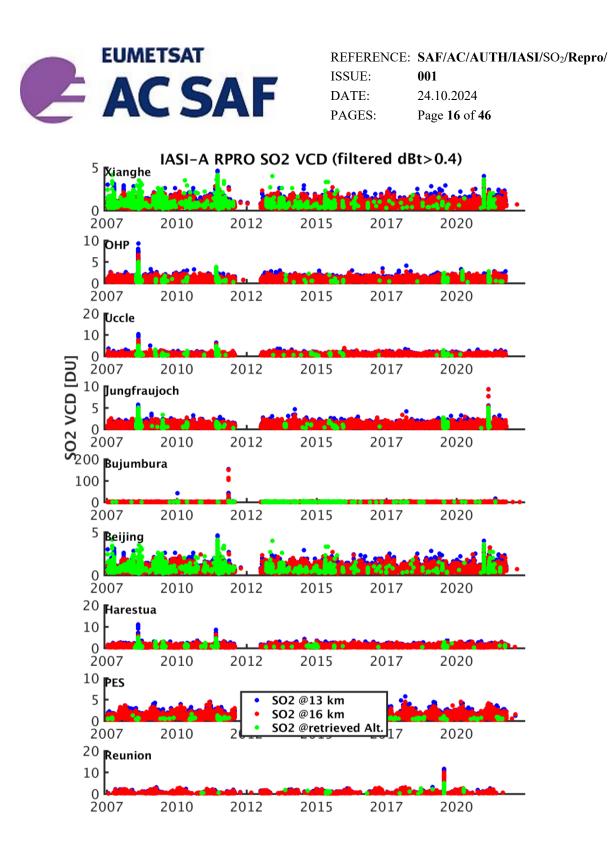
Several BIRA-IASB stations (http://uv-vis.aeronomie.be/groundbased/stations) cover part of the available datasets time-period (2007-2021, see Table 3) and the IASI SO<sub>2</sub> data have been extracted in a radius of 300km around these sites, as shown in Figure 6 for the IASI-A data and in Figure 7 for the IASI-B data.

As illustrated in Figure 6 for IASI-A, the overpasses have been calculated keeping all the valid points (i.e., points with  $\Delta Bt > 0.4K$ ) and then a stricter filtering applying the recommended  $\Delta Bt>1K$  threshold to only highlight the volcanic cases. As discussed in Section 2.2.1 a recommendation on using pixels with associated  $\Delta Bt > 0.4K$  if in the vicinity of pixels with associated  $\Delta Bt > 1K$  exists but is not implemented by default and is left to the users to apply it. We will test the impact of this choice in a few cases below.

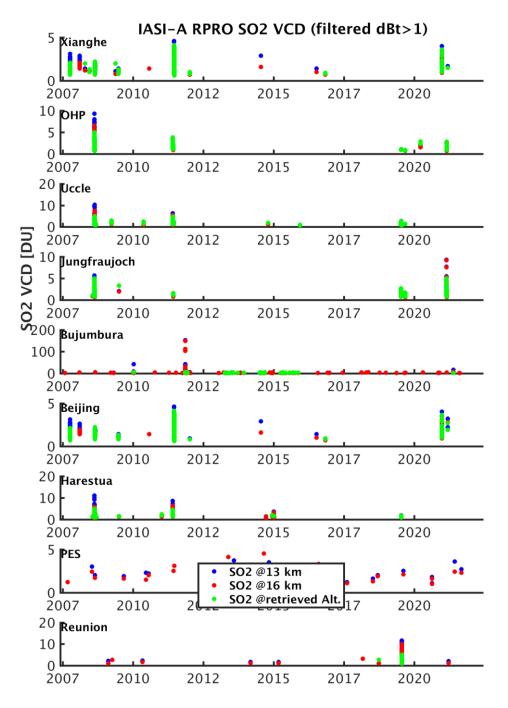
When applying the filter on  $\Delta Bt > 1K$ , several SO<sub>2</sub> plumes from large eruptions can be seen over several BIRA-IASB stations, with several plumes travelling around the Northern and/or Southern hemispheres for several days, such as:

- the Kasatochi eruption in August 2008 (Clarisse et al., 2012), which travelled around the northern hemisphere between 11/08/2008 and 06/09/2008;

- the Sarychev eruption on the 30/6/2009;
- the Eyjafjallajokull eruption on the 17/5/2010;
- the Nabro eruption between the 12 and 28 June 2011 (Clarisse et al., 2014);
- the Grimsvotn eruption at the end of May/early June 2011;
- the Nyamuragira eruptive period between 6/11/2011 and 6/12/2011;
- the Raikoke eruption in late June 2019;
- the Ubinas eruption on the 23/7/2019.

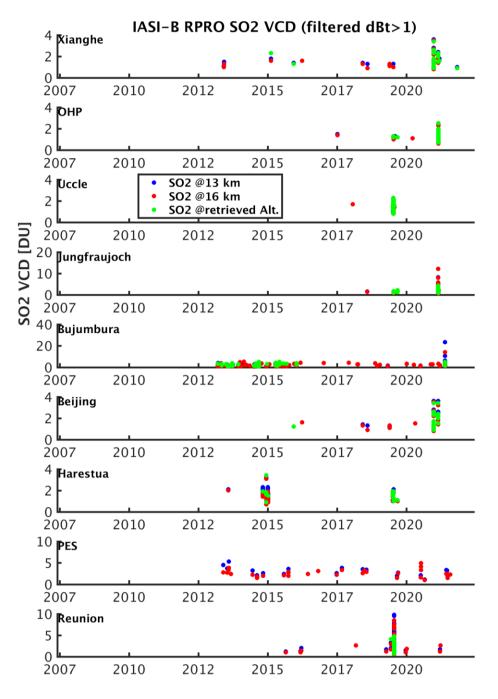






**Figure 6.**IASI-A SO<sub>2</sub> columns overpasses around the BIRA-IASB stations; SO<sub>2</sub> VCD estimates for 13km (blue dots), 16km (red dots) and for the retrieved height (green dots). Two filtering cases are shown, for  $\Delta Bt > 0.4K$  (top) and  $\Delta Bt > 1K$  (bottom).





**Figure 7.** IASI-B SO<sub>2</sub> columns overpasses around the BIRA-IASB stations for  $\Delta Bt > 1K$ ; SO<sub>2</sub> VCD estimates for 13km (blue dots), 16km (red dots) and for the retrieved height (green dots).



Station	Full measurement period	Information relevant for SO <sub>2</sub>
Uccle (50,5°N, 4°E, Belgium)	mini-MAXDOAS instrument from 5/2011 to 2/2016 (Gielen et al., 2014) and a new MAXDOAS instrument from 2/2017 to 2/2020 (Dimitropoulou et al 2020; 2022)	
<b>Jungfraujoch</b> (46.5°N, 7.98°E, <b>Switzerland</b> )	ZenithSky geometry since the '90 with a SAOZ, a MAXDOAS geometry from 7/2010 to 10/6/2019 (Hendrick et al, 2012) and a new Zenith (only VIS) since 3/2022	The MAXDOAS UV channel had permanent failure in mid- June 2019
<b>OHP</b> (44°N, 7.71°E, <b>France</b> )	MAXDOAS geometry from 2005 to March 2017 (Valks et al., 2011)	measuring in the spectral region from 323-390nm and is thus not covering the SO <sub>2</sub> analysis region (305-317.5nm)
<b>Xianghe</b> (39.75°N, 116.96°E, <b>China</b> )	MAXDOAS from 2010 to summer 2022 (Clémer et al., 2010; Hendrick et al, 2014)	instrumental problem between 7/2018 to 10/2019 (measuring only in zenith mode). The UV channel had permanent failure in end October 2021.
<b>Bujumbura</b> (3°S, 29°E, <b>Burundi</b> )	MAXDOAS from 11/2013 to mid 2017 (Gielen et al., 2017)	Problem with UV channel between 25/4/2014 and 14/4/2015 and since 16/8/2017.
<b>Reunion</b> (21°S, 55°E, <b>Reunion Island</b> )	MAXDOAS in LePort from 4/2016 to 1/2018 and on Maido since 6/2018	Problem with the degradation of the UV slit function since mid of 2021. Data not usable after.

Shown in the previous figures and considering the ground-based measurements limitations (as discussed in Table 3), five stations have at least one interesting eruption to look into in detail. Ground-based SO<sub>2</sub> analysis are not performed routinely, and only the time-periods with strong volcanic SO<sub>2</sub> signals revealed by both IASI and GOME2 observations are analysed.

The ground-based spectra have been analysed using the DOAS settings specified in Wang et al., 2014 (analysis in the 305-317.5 nm window) with a fixed reference spectrum close to the period of interest (on a day not affected by the eruption). The slant columns measured at zenith mode are used to



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estimate the stratospheric SO<sub>2</sub> content by using a geometrical stratospheric AMF that depends only on the solar zenith angle (Honninger et al., 2004):

$$AMF_{strato} = 1/cos(SZA)$$
(1)

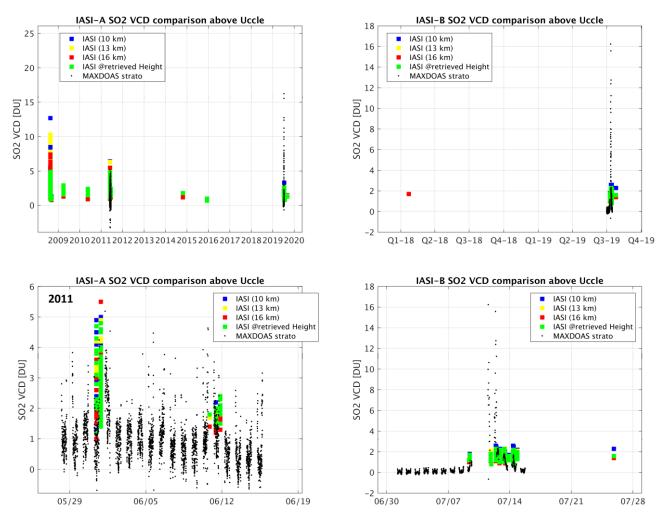
For large stratospheric eruptions, such as Kasatochi, Grimsvotn, Nabro, etc. (Carn et al., 2016 and Hopfner et al., 2015), both MAXDOAS and ZenithSky instruments are sensitive to the SO<sub>2</sub> in the stratosphere, and the stratospheric SO<sub>2</sub> content is obtained as:

$$VCD_{strato} = DSCD_{zenith} / AMF_{strato}$$
 (2)

In the following sub-sections, the different comparison cases for the stations listed in Table 3 are presented.

#### Uccle

An illustration of the IASI-A and IASI-B overpasses for Uccle is given in Figure 8 for the whole timeline on top and for two specific periods in 2011 for IASI-A and 2019 for IASI-B at the bottom.



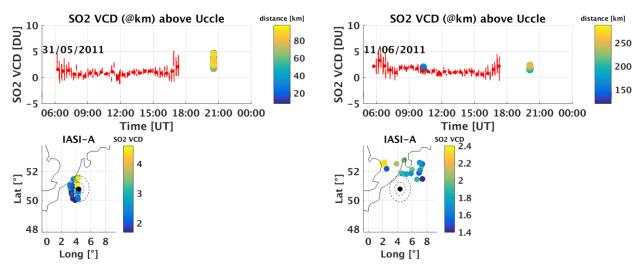


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**Figure 8.** Uccle SO<sub>2</sub> VCD time-series of the IASI-A and IASI-B overpasses (colored squares) and the analyzed ground-based data (black dots). Top: the whole time-series; Lower Left: June 2011 Grimsvotn eruption; Right: July 2019 Raikoke eruption.

The Uccle station can be challenging because the instrument measuring in the 2011-2016 period is a commercial mini-MAXDOAS system from Hoffmann (Gielen et al., 2014) and the SO<sub>2</sub> retrieval is quite noisy. From 2/2017 to 2/2020 a scientific grade MAXDOAS was measuring in Uccle, and the noise on the SO<sub>2</sub> is much reduced. We analysed the May-June 2011 and the July 2019 periods for the Grimsvotn and the Raikoke eruptions from these instruments.

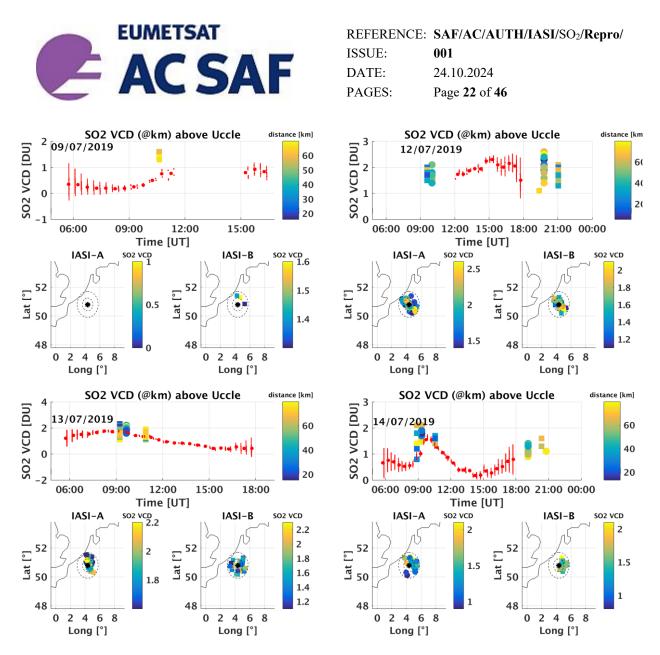
Two days in 2011 with IASI-A SO<sub>2</sub> collocations are shown in Figure 8 are illustrated in Figure 9 (31/5, left and 11/6, right). On the 31/5/2011 IASI-A only has pixels within 80km with  $\Delta$ Bt>1K in the nighttime overpass and that on the 11/06/2011 all the pixels are quite far away from the stations, at about 150 to 250km. The distance of the satellite pixel from the ground-station is shown as the colour of the satellite circles, with the ground-based assessment shown in red.



**Figure 9.** SO<sub>2</sub> VCD comparisons at Uccle for 2 days focusing on the Grimsvotn eruption in 2011, left and right columns. Each figure contains an upper panel with the SO<sub>2</sub> VCD time-series from the zenith measurements (in red) and from IASI (columns at the retrieved height, color-coded as a function of the station distance). The lower panel presents the geographical distribution of the IASI SO<sub>2</sub> VCD around the station (circles at 50km and 100km are plotted in dotted lines).

The few days with existing IASI-B SO<sub>2</sub> height seen in Figure 8 for 2019, are illustrated in Figure 10, keeping only the pixels that are within 80km of the Uccle site. IASI-A collocations are also available here. It can be seen that:

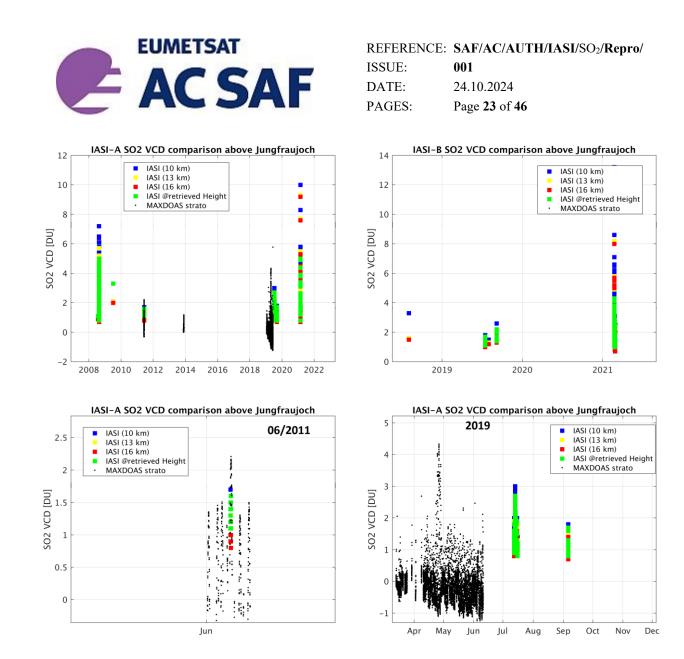
- On the first day (9/7/2019, show in the upper left), only IASI-B has some volcanic SO<sub>2</sub> pixels (with  $\Delta$ Bt>1K), and there is a tendency of increasing SO<sub>2</sub> during the morning.
- On the 12<sup>th</sup> of July 2019 (upper right) there are no temporal collocations, but IASI-A and IASI-B have volcanic signals in both the day and night overpasses, with similar SO<sub>2</sub> levels than the ground-based data, and with a tendency of larger columns in the night-time overpass, as for the growing SO<sub>2</sub> columns during the day.
- On the 13<sup>th</sup> and 14<sup>th</sup> July (bottom row), both IASI-A and IASI-B day overpasses are in agreement with the ground-based data, with values of about 2DU and with slightly larger values for IASI-A compared to IASI-B and a decreasing SO<sub>2</sub> load during the day.



**Figure 10.** SO<sub>2</sub> VCD comparisons at Uccle focusing on the July 2019 eruption (Raikoke). Days are left to right, top to bottom. Each panel presents an upper panel, with the SO<sub>2</sub> VCD time-series from the zenith measurements (in red) and from IASI (columns at the retrieved height, color-coded as a function of the station distance; circles for IASI-A, squares for IASI-B). The second panel presents the geographical distribution of the IASI-A (left) and IASI-B (right) SO<sub>2</sub> VCD around the station (circles at 50km and 100km are plotted in dotted lines).

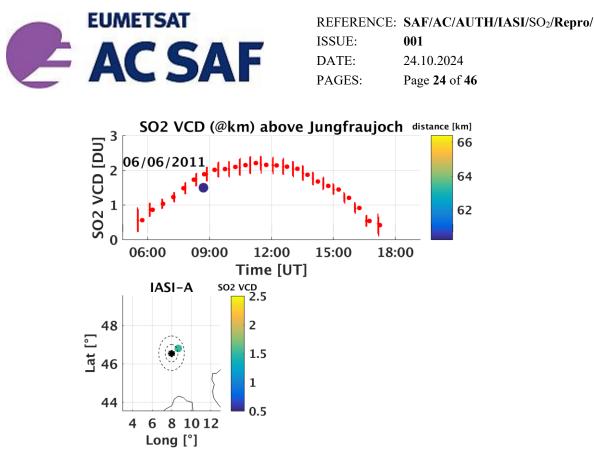
### Jungfraujoch

An illustration of the IASI-A and IASI-B overpasses for Jungfraujoch is given in Figure 11 for the whole time-series in the upper row and for June 2011 and in spring/summer 2019 when the Grimsvotn and the Raikoke plumes circulated around the Northern Hemisphere in the bottom row. As it can be seen, on the 6 of June 2011 there is an SO<sub>2</sub> enhancement for both ground-based and IASI-A, while unfortunately the MAXDOAS UV channel stopped measuring on 10<sup>th</sup> June 2019, so there is no possibility to compare for the 2019 case and later. During the Kasatochi eruption (August 2008) only the BIRA SAOZ was measuring (MAXDOAS installed in 2010), which is not measuring in the SO<sub>2</sub> region.



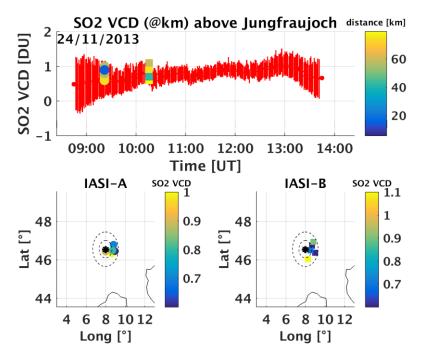
**Figure 11.** Jungfraujoch SO<sub>2</sub> VCD time-series of the IASI-A (left) and IASI-B (right) overpasses (colored squares) and the analyzed ground-based data (black dots). Up: the whole 2007-2021 time-series; Bottom left: zoom on June 2011 for the Grimsvotn eruption; Bottom right: zoom on summer 2019 for the Raikoke eruption.

The case of 6 June 2011 is shown in detail in Figure 12, with one IASI-A pixel with  $\Delta Bt>1$  K in the 80km around the site. The satellite SO<sub>2</sub> estimate is very close to the value observed from the ground for that day.



**Figure 12.** SO<sub>2</sub> VCD comparisons at Jungfraujoch on 6/6/2011 for the Grimsvotn eruption. The figure presents an upper panel, with the SO<sub>2</sub> VCD time-series from the zenith measurements (in red) and from IASI-A (SO<sub>2</sub> columns at the retrieved height, color-coded as a function of the station distance). The lower panels present the geographical distribution of the IASI-A SO<sub>2</sub> VCD around the station (circles at 50km and 100km are plotted in dotted lines).

In the [VR], the ground-based data at the end of 2013 revealed a volcanic plume that was also observed by the IASI-A SO<sub>2</sub> AERIS dataset, coming from an Etna plume from an eruption on the 23rd of November travelling north and reaching Jungfraujoch on the 24/11/2013. This eruption is not retained by the  $\Delta$ Bt>1K selection, but we can find it when looking at the IASI pixels with  $\Delta$ Bt>0.4K, as illustrated in Figure 13.





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**Figure 13.** SO<sub>2</sub> VCD comparisons at Jungfraujoch on 23/11/2013 for an Etna eruption. The figure presents an upper panel, with the SO<sub>2</sub> VCD time-series from the zenith measurements (in red) and from IASI-A (circles) and IASI-B (squares) (SO<sub>2</sub> columns at the retrieved height, color-coded as a function of the station distance, for all pixels with  $\Delta$ Bt>0.4K within 80km of the station). The lower panels present the geographical distribution of the IASI-A SO<sub>2</sub> VCD around the station (circles at 50km and 100km are plotted in dotted lines).

#### Bujumbura

An illustration of the IASI-A overpass for Bujumbura is given in Figure 14 for the whole time-series in the left plot and for the November 2011 Nyamuragira eruption in the right. It can be seen that the very high SO<sub>2</sub> VCD estimated for plume altitudes between 10 to 16km are in fact reduced to up to  $\sim$ 4DU when the retrieved SO<sub>2</sub> height is used as information for the SO<sub>2</sub> VCD estimation. The MAXDOAS data only measured during the period 11/2013 to mid-2017, so this big eruption is not covered.

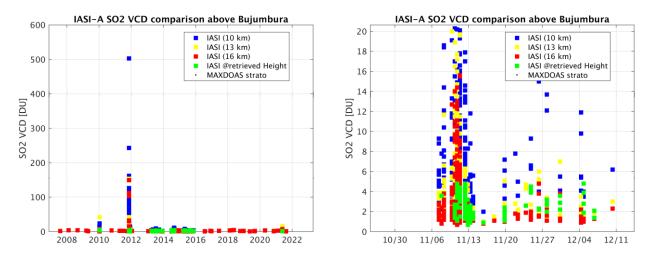
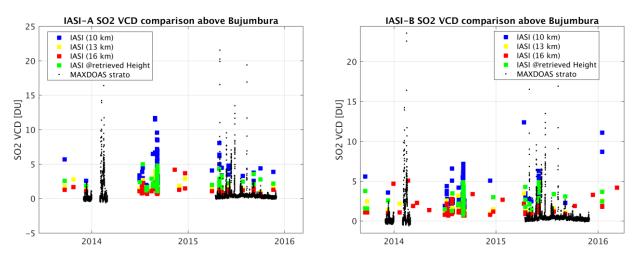


Figure 14. Bujumbura SO<sub>2</sub> VCD time-series of the IASI overpasses (colored squares). Left: the whole 2007-2021 time-series; right: zoom on Nyamuragira eruption of 6/11 to 6/12/2011.



**Figure 15.** Bujumbura SO<sub>2</sub> VCD time-series of the 2013-2016 IASI-A (left) and IASI-B (right) overpasses (colored squares and the available retrieved ground-based data (black dots).



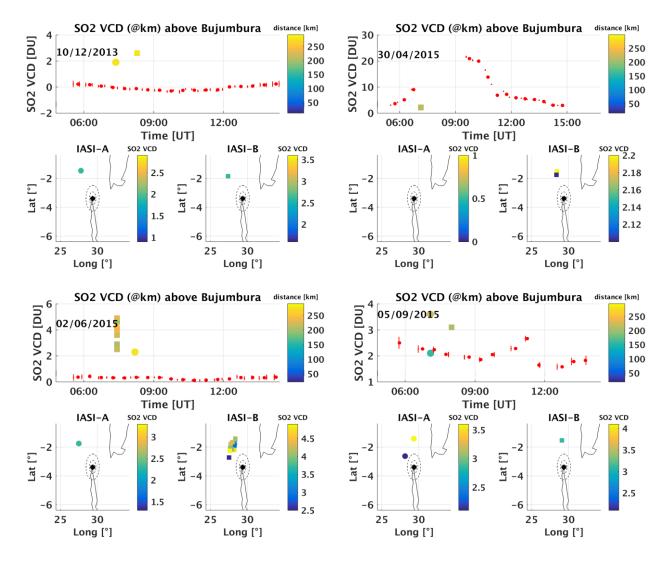
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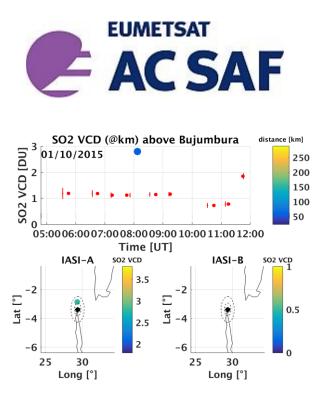
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The ground-based data were analysed for the 2013-2015 period, and data are presented in Figure 15. Unfortunately, there are several interruptions due to problems with the UV channel of the instrument (for e.g. in 2014, the data stops on the 26<sup>th</sup> of April and only starts again in April 2015) and the measurements stop on August 2017. Several SO<sub>2</sub> peaks are seen both in the IASI overpasses and in the ground-based signal, but the direct comparison is difficult because many peaks are not on the same days, probably due to some tropospheric SO<sub>2</sub> contamination of the ground-based dataset, and on the few common days (10 Dec-2013, 30-Apr-2015, 02-Jun-2015, 05-Sep-2015 and 01-Oct-2015) the IASI pixels with  $\Delta$ Bt >1K are quite distant from the station (up to 300km), as shown in Figure 16.

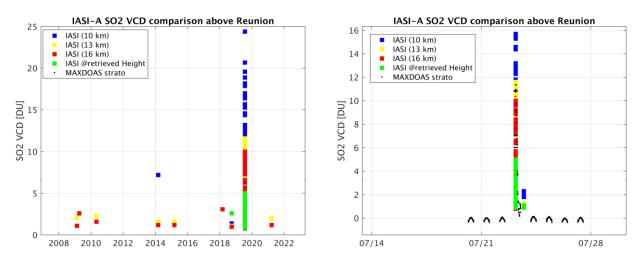




**Figure 16.** SO<sub>2</sub> VCD comparisons at Bujumbura for different eruptive days. The figure presents an upper panel, with the SO<sub>2</sub> VCD time-series from the zenith measurements (in red) and from IASI-A (circles) and IASI-B (squares) (SO<sub>2</sub> columns at the retrieved height, color-coded as a function of the station distance). The lower panels present the geographical distribution of the IASI-A (left) and IASI-B (right) SO<sub>2</sub> VCD around the station (circles at 50km and 100km are plotted in dotted lines).

#### Reunion

An illustration of the IASI-A overpass for Reunion is given in Figure 17 for the whole time-series (left) and for 23th of July 2019, when the Ubinas eruption plume passed close to the Reunion island in the right (https://volcano.si.edu/showreport.cfm?doi=GVP.WVAR20190724-354020).



**Figure 17.** Reunion  $SO_2$  VCD time-series of the IASI overpasses (colored squares) and the analyzed ground-based data (black dots). Left: the whole 2007-2021 time-series; Right: zoom on July 2019 for the Ubinas plume on the 23th.

BIRA-IASB measured with a MAXDOAS in LePort (on the coast) between April 2016 and January 2018, and then moved the instrument to the Maido site (at 2360m) and has been measuring there since June 2018. The data from the Maido site are thus compared to the SO<sub>2</sub> at the retrieved altitude for the IASI-A and IASI-B overpasses in Figure 18. It can be seen that on the 23/07/2019, during the day the

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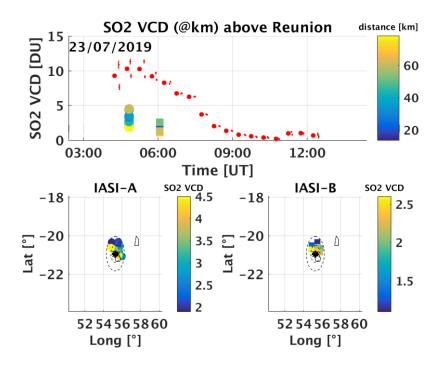
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SO<sub>2</sub> from the ground-based decrease from about 10DU in the morning to about zero later in the day (from 9h UT). Similarly, the SO<sub>2</sub> values decrease from the IASI-A to the IASI-B overpass time (around half an hour later).



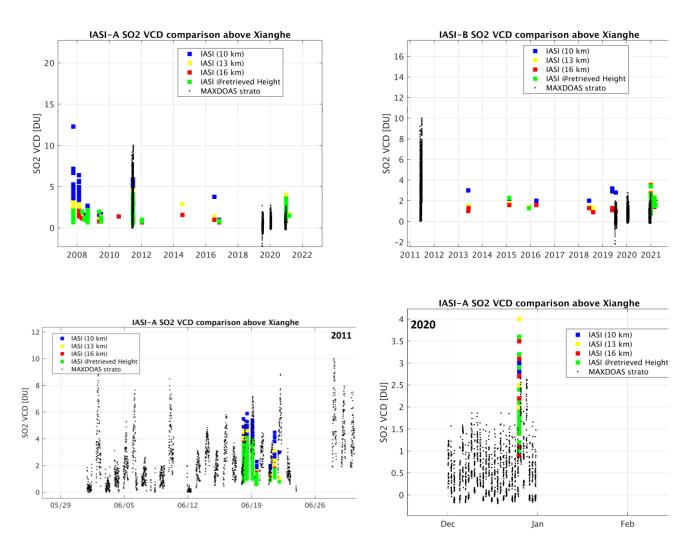
**Figure 18.** SO<sub>2</sub> VCD comparisons at Reunion focusing on the Ubinas eruption on 23/07/2019. The figure presents an upper panel, with the SO<sub>2</sub> VCD time-series from the zenith measurements (in red) and from IASI-A (circles) and IASI-B (squares) (SO<sub>2</sub> columns at the retrieved height, color-coded as a function of the station distance). The lower panels present the geographical distribution of the IASI-A (left) and IASI-B (right) SO<sub>2</sub> VCD around the station (circles at 50km and 100km are plotted in dotted lines).

### Xianghe

Figure 19 present the time-series of the full IASI-A and IASI-B overpasses (upper row), and for June 2011 (first eruption after the installation of the MAXDOAS in 2010) during when the Nabro plume passed around the Xianghe station in the bottom left as well as in December 2020 in the bottom right.



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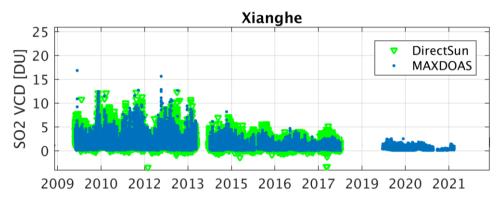
**Figure 19.** Xianghe SO<sub>2</sub> VCD time-series of the IASI-A and IASI-B overpasses (colored squares) and the analyzed ground-based data (black dots). Top row: the whole time-series; bottom: June 2011 Nabro eruption (left) and December 2020 (right).

As can be seen in Figure 19 for June 2011 and as discussed in past validation exercises [VR], the Xianghe station is more challenging because of the SO<sub>2</sub> pollution that is observed by the MAXDOAS instrument all year long (Wang et al., 2014), especially before 2018. This needs to be disentangled from the volcanic signal. To that end, an additional approach was tested, using also the direct sun and the MAXDOAS modes and comparing the results obtained in the different geometries. The direct sun total SO<sub>2</sub> VCD and the MAXDOAS tropospheric SO<sub>2</sub> columns from the profiling algorithm (Wang et al. 2014) are inter-compared (see Figure 20) and to the SO<sub>2</sub> VCD obtained using stratospheric VCD estimated from the zenith measurements using formula (2) as discussed before and performed for the other stations (see Figure 21 for the June 2011 period). Most of the time, the total SO<sub>2</sub> from the direct sun signal is larger, sign of some additional (stratospheric) SO<sub>2</sub> content. The difference between the direct sun and the MAXDOAS data are compared in the lower plot of Figure 21. For days where both datasets agree well, we are confident of the retrieved stratospheric SO<sub>2</sub>, and this can be compared to IASI SO<sub>2</sub> data, as shown in Figure 22. The IASI plume height for the Nabro eruption has been estimated from

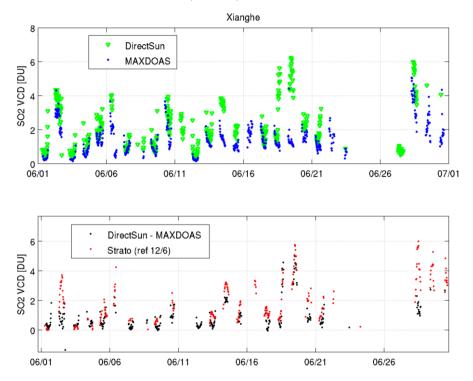


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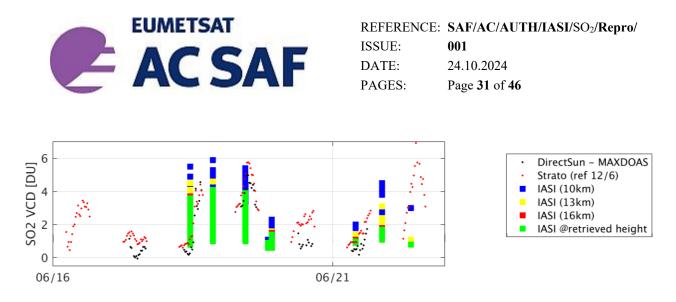
IASI in Clarisse et al., 2014 and is of about 16km around the Xianghe station on the 19 June 2011 (Clarisse et al., 2014, their Figure 10). The SO<sub>2</sub> at the retrieved height is shown in Figure 22 with also the SO<sub>2</sub> values estimated for heights from 10 to 16km, and it compare well with the SO<sub>2</sub> at 16km height, as expected. The IASI-A SO<sub>2</sub> columns at the retrieved height are thus compared day by day in Figure 23 to the ground-based data. When data with  $\Delta$ Bt>1K exist, only the pixels within a radius of 80km around Xianghe are presented (18 and 19 June 2011), otherwise the pixels within 300km are shown.



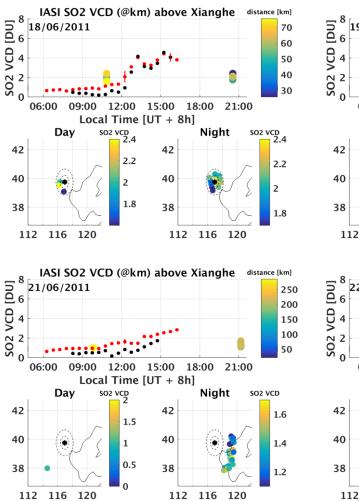
**Figure 20.** SO<sub>2</sub> VCD comparisons at Xianghe for total SO<sub>2</sub> from direct sun geometry (in green) and tropospheric SO<sub>2</sub> VCD from MAXDOAS (in blue).

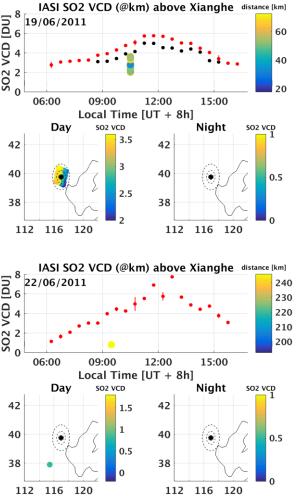


**Figure 21.** SO<sub>2</sub> VCD comparisons at Xianghe focusing on June 2011. The first panel shows total SO<sub>2</sub> from direct sun geometry (in green) and tropospheric SO<sub>2</sub> VCD from MAXDOAS (in blue). The second panel shows stratospheric SO<sub>2</sub>, as retrieved from the total minus tropospheric VCD (in black) and as retrieved using a stratospheric AMF from zenith measurements (in red).



**Figure 22.** SO<sub>2</sub> VCD comparisons at Xianghe focusing on June 2011 around the Nabro plume overpass. IASI columns in a radius of 300km around the stations are reported for different heights and the stratospheric VCD are estimated from the total direct sun minus tropospheric MAXDOAS VCD (in black) and from zenith measurements with a fixed DOAS reference taken on 12/06/2011 at noon (in red).







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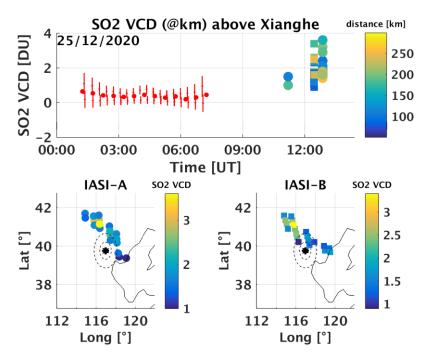
**Figure 23.** SO<sub>2</sub> VCD comparisons at Xianghe focusing on the 18 to 22 June 2011. Days are left to right, top to bottom. Each day present an upper panel, with the SO<sub>2</sub> VCD time-series reported from the total direct sun minus tropospheric MAXDOAS (in black), from the zenith measurements (in red) and from IASI-A (columns at the retrieved height, color-coded as a function of the station distance). The second panel presents the geographical distribution of the IASI SO<sub>2</sub> VCD around the Xianghe station (circles at 50km and 100km are plotted in dotted lines) for the day (left) and the night (right) overpasses.

From Figure 23, it can be seen that:

- the ground-based signal increases in the afternoon of 18/6/2011, from 2 to 4DU, comparable with the 2 to 2.5DU seen by IASI-A in the day and night overpasses;
- a stable to increasing signal on 19/6/2011 with 3 to 6 DU is seen by the ground-based, while 2 to 4 DU are seen by IASI during the morning overpass, with larger values for pixels west of the station;
- on the following days the SO<sub>2</sub> signal is reduced: no IASI-A signal on the 20 June and values are between 1 and 2 DU on 21 and 22 June, with a small number of pixels far away the station. From the ground, a signal between 1 and 2 DU is also seen on the 21 June, while on the 22 the signal is larger, but only data from the zenith measurements could be retrieved, so we trust less this day.

There are thus two good days to compare with, with two others have pixels too distant from the station. In this case, permitting also pixels with  $\Delta Bt$  between 0.4 and 1 (and not only >1K) increases the number of IASI pixels to compare with, however also increases the noise.

The volcanic signal in December 2020 is seen by both IASI-A and IASI-B, with some increase in the ground-based stratospheric VCD data (the direct sun mode was not working any more), but unfortunately, for both satellite instruments we only have a night overpass, as can be seen in Figure 24.





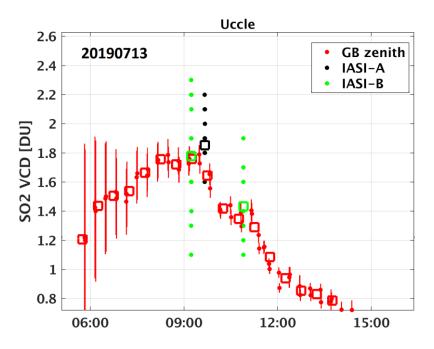
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**Figure 24.** SO<sub>2</sub> VCD comparisons at Xianghe focusing on the 25 December 2020. In the upper panel, the SO<sub>2</sub> VCD time-series are reported from the zenith measurements (in red) and from IASI-A (dots) and IASI-B (squares) (SO<sub>2</sub> columns at the retrieved height, color-coded as a function of the station distance). The lower panels present the geographical distribution of the IASI-A (left) and IASI-B (right) SO<sub>2</sub> VCD around the Xianghe station (circles at 50km and 100km are plotted in dotted lines) for the day (left) and the night (right) overpasses.

#### 3.2.1 Conclusions from the comparisons with the BIRA-IASB ground-based stations

Although the relatively small number of valid coincident cases (only 8 IASI-A and 5 IASI-B comparison days within pixels within 80km of the stations), qualitatively there seems to be a good agreement, with ground-based stations generally showing enhanced SO<sub>2</sub> signals just before/during/after the enhanced IASI SO<sub>2</sub> overpasses and the order of magnitude of the SO<sub>2</sub> signals are generally comparable. IASI-A and IASI-B usually display SO<sub>2</sub> temporal variations coherent with what seen from the ground.

To look at the results in a more quantitative way, we can average the ground based SO<sub>2</sub> data for half an hour around the IASI overpass time ( $\pm 15$ min) and compare them to the mean SO<sub>2</sub> VCD loads from the selected IASI pixels (see Figure 25 for an illustration). We can do this exercise for the different IASI SO<sub>2</sub> VCD estimates, based on the fixed 7 to 25km heights and on the retrieved altitude. The resulting mean SO<sub>2</sub> VCD are summarized in Table 4 and Table 5 for IASI-A and IASI-B respectively. Please note that due to the  $\pm 15$ min temporal coincidence restriction, some of the cases presented above are not retained here (e.g., Bujumbura on 1/10/2015).



**Figure 25.** Illustration of the quantitative comparison method for Uccle on 13/07/2019, taking 30-minutes averages (squares) from the original ground-based and IASI datasets.



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**Table 4.** Mean SO<sub>2</sub> VCD values for the ground-based and IASI-A different options, for the good condition cases (pixels from daily overpass within 80km of the station).

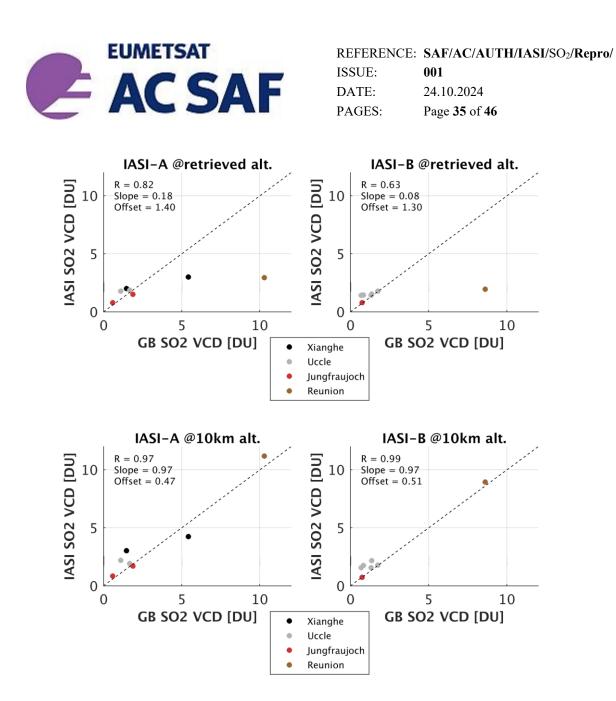
IASI-A				SO <sub>2</sub> VCD [DU]						
year	month	day	Station	GB	retrieved height	7km	10km	13km	16km	25km
2011	6	18	Xianghe	1,48	2,00	7,10	3,03	2,53	2,13	2,43
2011	6	19	Xianghe	5,43	2,99	8,52	4,24	3,10	2,96	3,43
2019	7	13	Uccle	1,66	1,85	4,26	1,92	1,90	1,62	1,78
2019	7	14	Uccle	1,09	1,78	3,33	2,20	1,84	1,61	1,68
2011	6	6	JFJ	1,88	1,50	n/a	1,70	1,40	1,20	1,40
2013	11	24	JFJ	0,59	0,80	1,78	0,83	0,83	0,82	0,70
2019	7	23	Reunion	10,28	2,95	20,62	11,19	5,57	4,87	6,47

**Table 5.** Mean SO<sub>2</sub> VCD values for the ground-based and IASI-B different options, for the good condition cases (pixels from daily overpass within 80km of the station).

IASI-B				SO <sub>2</sub> VCD [DU]						
year	month	day	Station	GB	retrieved height	7km	10km	13km	16km	25km
2019	7	9	Uccle	0,69	1,43	4,50	1,57	1,40	1,30	1,43
2019	7	13	Uccle	1,76	1,78	4,48	1,79	1,76	1,48	1,63
2019	7	13	Uccle	1,33	1,43	4,53	1,57	1,46	1,26	1,41
2019	7	14	Uccle	0,84	1,46	3,14	1,76	1,38	1,16	1,28
2019	7	14	Uccle	1,37	1,54	3,66	2,16	1,60	1,40	1,55
2013	11	24	JFJ	0,74	0,78	1,86	0,72	0,82	0,74	0,66
2019	7	23	Reunion	8,61	1,94	16,67	8,94	4,52	3,81	4,92

Generally, the results are very similar for IASI-A and IASI-B, with a good agreement of the groundbased data with the IASI SO<sub>2</sub> VCD at the retrieved altitude for moderate SO<sub>2</sub> content (cases with SO<sub>2</sub> VCD below 5 DU). Larger differences appear for the Reunion 23/7/2019 case and the Xianghe 19/6/2011 case, which have ground-based columns around 10 and 5 DU respectively, while IASI retrieve smaller SO<sub>2</sub> values when considering the retrieved altitude. The SO<sub>2</sub> estimate for 10km height appears closer to the ground-based values.

The summary result comparisons for all the cases together are presented in Figure 26 for the IASI  $SO_2$  VCD at the retrieved altitude and for the 10km altitude assumption. Correlation coefficients are very good (about 0.82/0.63 for IASI-A and IASI-B and up to 0.97/0.99 for the 10km assumption) with linear regression slope close to one in the second case.



**Figure 26.** SO<sub>2</sub> VCD comparisons summary for the different IASI SO<sub>2</sub> estimates: SO<sub>2</sub> retrieved height (upper) and fixed 10km height (lower) for IASI-A (left) and IASI-B (right).



## 4. SULPHUR DIOXIDE ALTITUDE VALIDATION

Accurate determination of the location, height and loading of SO<sub>2</sub> plumes emitted by volcanic eruptions is essential for aviation safety. The SO<sub>2</sub> layer height is furthermore one of the most critical parameters that determine the impact on the climate. The height of volcanic ash columns are often estimated by local observers with mostly unknown accuracy. The plume height can also be determined using aircraft, ground-based radar or LIDAR but such observations are often not available and many volcanic eruptions in remote areas remain not observed. In addition, volcanic plumes containing SO<sub>2</sub> but not ash cannot be seen directly by either ground-based or space-born lidars.

 $SO_2$  in the atmosphere has important impacts on chemistry and climate at both local and global levels. Natural sources account for ~30% of  $SO_2$  emissions. Next to contributions from volcanic activity, these include emissions from marine phytoplankton and a small contribution from soil and vegetation decay. However, by far the largest contributions in global  $SO_2$  production are from anthropogenic sources. These account for the remaining 70% of global emissions and primarily relate to fossil fuel burning, with smaller contributions from smelting and biomass burning.

#### 4.1.1 S5P/TROPOMI SO<sub>2</sub> Layer Height

While satellite instruments, in principle, provide global products e.g. from SEVIRI (Second Generation Spin-stabilised Enhanced Visible and Infra-Red Imager) or AIRS (Atmospheric Infra-Red Sounder), they have no or little vertical resolution. Although retrievals of SO2 plume height have been carried out using satellite UV backscatter measurements from e.g. OMI (Ozone Monitoring Instrument) or GOME2, until now such algorithms are up to now very time-consuming, since the spectral information content and its characterization require computationally demanding radiative transfer modelling. Due to the high spatial resolution of TROPOMI (Tropospheric Ozone Measurement Instrument) aboard S5P (Sentinel-5p) and consequent large amount of data, an SO<sub>2</sub> layer height algorithm has to be very fast. The Sentinel-5P Innovation project (S5p+I) has been initiated to develop novel scientific and operational applications, products and retrieval methods that exploit the potential of the Sentinel-5p mission's capabilities beyond its primary objective. The SO<sub>2</sub> Layer Height (SO<sub>2</sub>LH) theme was dedicated to the generation of an SO<sub>2</sub> layer height product for Sentinel-5p taking into account data production timeliness requirements. In this report, we will compare the ESA S5P+I SO<sub>2</sub> Layer Height product, which is based on the algorithm of Efremenko et al., 2017, further improved by Hedelt et al., 2019, validated by Koukouli et al., 2022 and assimilated successfully into the CAMS forecasting system, as demonstrated by Inness et al., 2022.

The S5P SO<sub>2</sub> LH data used in this report are available upon request from the ESA S5P+I: SO2LH project Data Pool<sup>2</sup> where the final version of the SO<sub>2</sub> LH data, v4.0, is available. We should note here that this algorithm has been implemented during 2023 into the operational ESA MPC-ATM S5P dataset chain and is part of the operational S5P SO<sub>2</sub> dataset since July 2023<sup>3</sup>. Due to restrictions of the algorithm, discussed in Hedelt et al., 2019 and Koukouli et al., 2022, a number of flags are applied to the dataset for quality assured results. The main one affecting this analysis is about the retrieved SO<sub>2</sub> load by the algorithm which is set to be > 20 D.U. As a result, shown in the section below, a number of the thinner SO<sub>2</sub> plumes observed by the IASI instruments which have no such restriction, are missing from the S5P SO<sub>2</sub> LH data.

<sup>&</sup>lt;sup>2</sup> Data pool | Sentinel-5p Innovation (dlr.de)

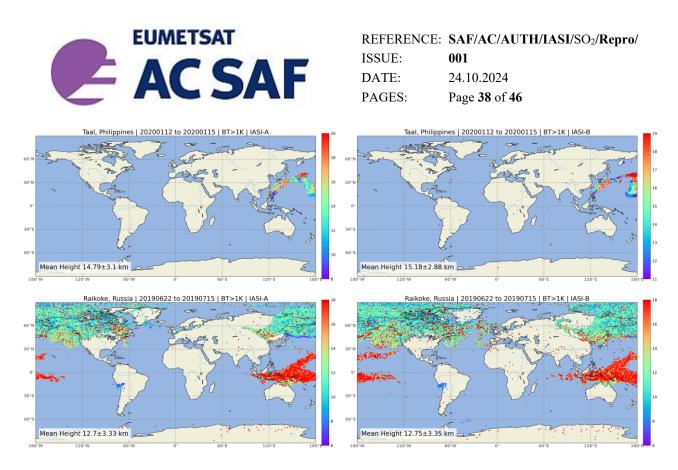
<sup>&</sup>lt;sup>3</sup> <u>Sentinel-5 Precursor Level 2 Sulphur Dioxide - Sentinel Online (copernicus.eu)</u>



#### 4.1.2 Volcanic eruptions studied

For the requirements of this report, five major eruptions during the common timeline of S5P and IASI are studied. In detail:

- On 22 June 2019 at 04:00h local time, the Raikoke stratovolcano located on the Kuril Islands (Russia, 551 m summit elevation) erupted explosively (VEI ≥ 4) after being dormant since 1924. There were several strong distinct explosions, producing a dense ash and SO<sub>2</sub> plume that rose until 13 km altitude the first days and was entrained into the stratosphere (see Sennert, 2019). This was the strongest volcanic eruption since the Merapi eruption in 2011, producing a colossal SO<sub>2</sub> plume with an SO<sub>2</sub> loading of more than 900 DU on 22 June 2019, that was dispersed by strong winds over Russia and North America, and was detectable even two weeks after the volcanic eruption. MLS was able to perform SO<sub>2</sub> profile measurements and also CALIPSO was able to detect the ash plume during the first days after the eruption.
- The **Taal** Volcano in Batangas, Philippines erupted on the afternoon of **January 12<sup>th</sup>**, **2020**, 43 years after its previous eruption in 1977. Stronger explosions began around 3 pm that spewed an ash column exceeding a kilometre high. By 7:30 pm, volcanic activities intensified as continuous eruptions generated a tall 10 to 15 kilometres steam-laden tephra column. On January 13, the activity on Taal's main crater had transitioned into a lava fountain.
- Nishinoshima lies about 600 miles (1,000 km) south of Tokyo, Japan, in the Ogasawara Arc. A young volcano, it has gone through several expansions in recent history. During a major eruption in 1973-1974, several new islands coalesced and expanded the size of the small island. In 2013, another vigorous eruption, which began offshore and continued until late 2015, eventually covered the island with lava flows and again enlarged the island. The volcano has been increasingly active through most of 2020, with July 2020 seeing extreme activity, including a series of powerful vulcanian-type eruptions from July 30-August 1st, characterized by dense, ash-rich cloud of gas exploding from the volcanic crater and typically rising high above the peak.
- On the morning of **April 9<sup>th</sup> 2021**, **La Soufrière** volcano, on the Caribbean island of Saint Vincent, began erupting, spewing ash at least 25,000 feet in the air. The volcano continued to erupt over the next several days, with multiple violent explosions. Ash blanketed Saint Vincent and winds carried ash to Barbados, about 120 miles east. A violent eruption on April 12<sup>th</sup> generated pyroclastic flows, a high-density mix of hot lava blocks, pumice, ash, and volcanic gas that moves at very high speeds down volcanic slopes. La Soufrière last erupted in 1979.
- Mt Etna, in Sicily, is a continuous outgassing volcano and one of the tallest active volcanoes in Europe. Beginning in February 2021, Mount Etna began a series of explosive eruptions, which have had an impact on nearby villages and cities, with volcanic ash and rock falling as far away as Catania. As of 12 March 2021, the volcano has erupted 11 times in three weeks. The eruptions have consistently sent ash clouds over 10 km (33,000 ft) into the air, closing Sicilian airports.

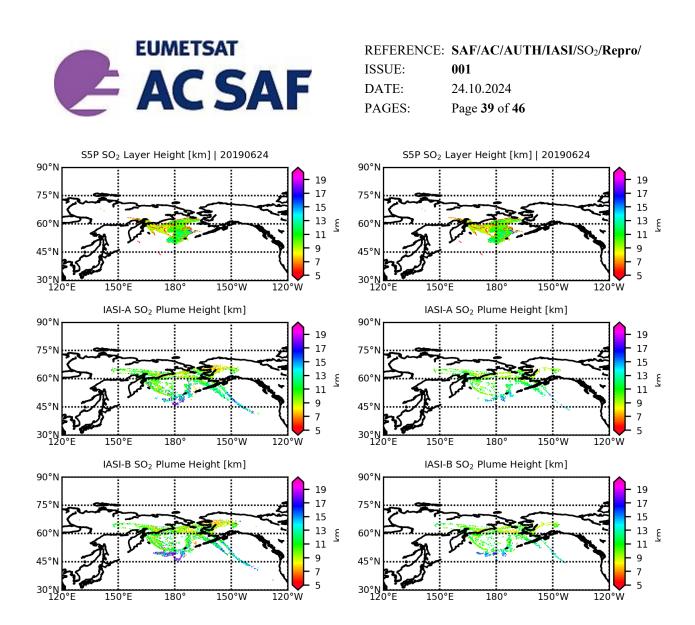


**Figure 27.** Examples of the IASI-A [left] and IASI-B [right]  $\Delta Bt > 1K$  SO<sub>2</sub> plume height during the Taal 2020 eruption [upper] and the Raikoke 2019 eruption [lower]. Note that during Raikoke, there was a synchronous volcanic eruption in Indonesia.

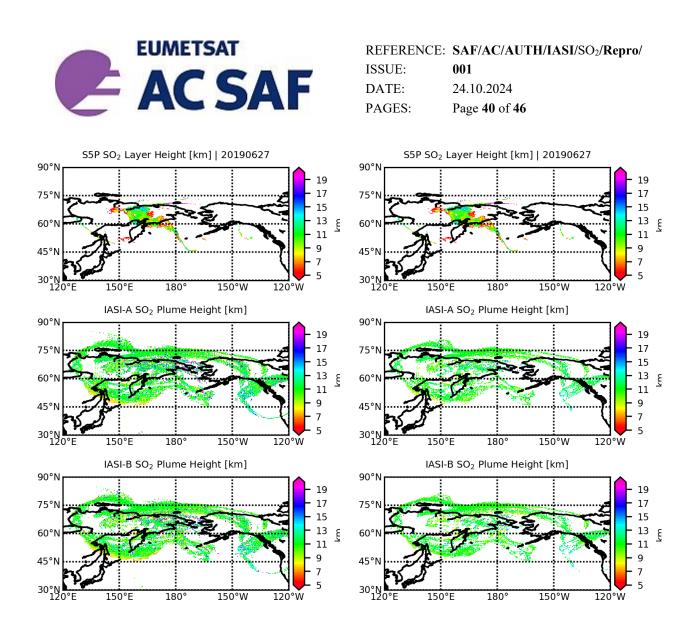
#### 4.1.3 Comparisons with the S5P/TROPOMI SO<sub>2</sub> layer height

In Figure 28 and Figure 29 we present two example days of the Raikoke eruption for demonstrational purposes. In the former the second day of the eruption is shown, the 24<sup>th</sup> of June 2019, while in the latter, the fifth day, the 27<sup>th</sup> of June 2019. In the upper panels the S5P SO<sub>2</sub> layer heights are presented, while IASI-A and IASI-B follow. The left column contains IASI data with  $\Delta Bt > 0.4$  K and the right  $\Delta Bt > 1$ K. Two main features are revealed by these pictorial investigations:

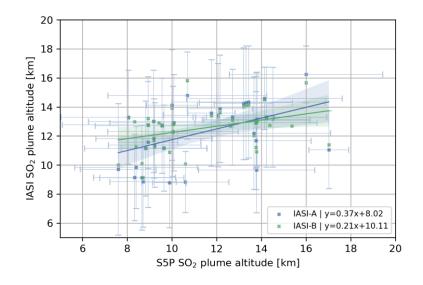
- The thickest, central, parts of the SO<sub>2</sub> plume appear to be missing from the IASI SO<sub>2</sub> plume heights shown in Figure 28. This part contains the highest SO<sub>2</sub> loads, within one day from the main eruption. These pixels are associated with forecast pressure and temperature profiles in the retrieval analysis and have been filtered by the use of the relevant so2\_qflag. Adding those pixels to the analysis does not alter the comparisons between collocated TROPOMI and IASI SO<sub>2</sub> layer heights in any significant way and in the worst case reduces the daily mean IASI SO<sub>2</sub> layer height by < 500m, well within the standard deviation levels (see Appendix.) Furthermore, S5P is unable to sense the thinner filament-like plumes revealed by IASI due to the minimum SO<sub>2</sub> load constraint imposed by the algorithm, as discussed above.
- This issue is even more pronounced in Figure 29 where IASI reports on the spread of the volcanic plume over a wide geographical range, which is not picked up by S5P.



**Figure 28.** The second day of the Raikoke eruption, on the  $24^{th}$  of June 2019, is shown for the S5P SO<sub>2</sub> layer height [upper], the IASI-A SO<sub>2</sub> plume height [middle] and the IASI-B SO<sub>2</sub> plume height [lower] for IASI data with  $\Delta Bt > 0.4$  K [left column] and  $\Delta Bt > 1$ K [right column.]



**Figure 29.** The fifth day of the Raikoke eruption, on the 27<sup>th</sup> of June 2019, is shown for the S5P SO<sub>2</sub> layer height [upper], the IASI-A SO<sub>2</sub> plume height [middle] and the IASI-B SO<sub>2</sub> plume height [lower] for IASI data with  $\Delta$ Bt > 0.4 K [left column] and  $\Delta$ Bt > 1K [right column.]



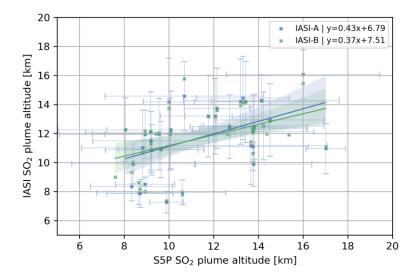


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**Figure 30.** Scatter plot of the daily mean SO<sub>2</sub> heights sensed by S5P, IASI-A [blue] and IASI-B [green] for all eruptive periods enumerated in **Table 6**. Upper plot for IASI data with  $\Delta Bt > 0.4$  K and lower plot for IASI data with  $\Delta Bt > 1.0$  K.

In Figure 30, scatter plots of the daily mean SO<sub>2</sub> heights sensed by S5P compared separately to IASI-A and IASI-B are presented. The stricter flagging of the IASI data with  $\Delta Bt > 1.0$  K improves both the slope and y-intercept of the comparisons, shown insert in the figures. In detail, for the comparison to IASI-A with  $\Delta Bt > 1.0$  K, the S5P SO<sub>2</sub> heights range between 8.06 and 17.02km, while IASI-A range between 7.27 and 16.07km for the 37 collocation days with a correlation coefficient of 0.50. For the comparison to IASI-B with  $\Delta Bt > 1.0$  K, the S5P SO<sub>2</sub> heights range between 7.6 and 17.02km, while IASI-B range between 7.36 and 15.76km for the 43 collocation days and a correlation coefficient of 0.47.

When comparing with IASI-A, the mean S5P SO<sub>2</sub> height is  $11.5\pm2.4$  km and IASI-A  $11.8\pm2.1$ km, while for IASI-B the mean heights are  $11.5\pm2.5$ km and  $11.7\pm2.0$ km, respectively. We overall report that the IASI-A and IASI-B SO<sub>2</sub> heights reside within 2.5km from the S5P SO<sub>2</sub> height estimates, well within the currently acceptable ranges for this product.



## 5. SUMMARY AND RECOMMENDATIONS

The dual aim of this validation report is to assert the capability of the EUMETSAT operational algorithm v6.6 to accurately retrieve IASI/MetopA and /MetopB total SO<sub>2</sub> columns and SO<sub>2</sub> plume heights on a global scale. For that end, Climate Data Recods (CDR) for IASI/MetopA between 01 Jan 2008 and 15 Oct 2021, and for IASI/MetopB between 20 Feb 2013 and 31 Dec 2021, were delivered by EUMETSAT to the validation teams.

In terms of the CDR netcdf files *per se*, we recommend that EUMETSAT considers in future releases to follow a more generic native-to-ncdf converter. Currently, differences in analysis were identified by the validation teams [not shown in this report] depending on which python module was employed in reading the CDR netcdf files. We further suggest the EUMETSAT considers releasing the IASI SO<sub>2</sub> CDR product in a harp-compliant netcdf format<sup>4</sup> in the future which would greatly increase the ease-of-use by the typical user.

In terms of assessing the suggested filtering on the SO<sub>2</sub> columns (section 2.2.1), we found that a number of pixels with high reported SO<sub>2</sub> columns are excluded when limiting the dataset to a  $\Delta$ Bt> 1K. Furthermore, we find the recommendation provided in the [PUM, section 5.2.2] that users *only look at the retrievals in the neighbourhood of* so<sub>2</sub>*bt\_difference* > 1K pixels, and not use the pixels with a so<sub>2</sub>*bt\_difference* < 0.4K (not enough SO<sub>2</sub> to have a reliable retrieval). The size of the region around so<sub>2</sub>*bt\_difference* > 1K pixels is let to the user's estimate. as non-viable for the typical user. We recommend that a volcanic flag is introduced in the dataset which will have a pre-defined region as "valid" which would still permit the expert user in making their own relevant calculations as they see fit for their work.

In terms of assessing the inter-sensor consistency (section 2.2.2) we report that both IASI-A ad IASI-B provide similar enough volcanic SO<sub>2</sub> retrieved columns which would enable the users to merge the two datasets and create a larger sample, always considering the temporal difference in their respective orbits.

In terms of comparisons against the S5P/TROPOMI volcanic SO<sub>2</sub> columns at an assumed plume height of 7km, section 3.1, we report that IASI-A and IASI-B nicely depict the major volcanic eruptions studied in this work, in terms of the location and spread of the plume on the day of, and following, the eruption. The overall levels of the IASI SO<sub>2</sub> retrieved columns are shown to be higher than those reported by S5P/TROPOMI both as SO<sub>2</sub> loads as well as SO<sub>2</sub> masses per day. Overall, we conclude that IASI-A and -B can be used in volcanic eruption monitoring and assessment.

In terms of comparisons against the BIRA MAX-DOAS ground-based observations, for the 8 IASI-A and 5 IASI-B temporally collocated days with pixels within 80km of the stations, a satisfactory agreement was found with correlation coefficients of about 0.82 for IASI-A and 0.63 for IASI-B for the retrieved SO<sub>2</sub> load and up to 0.97/0.99 for the SO<sub>2</sub> load with a 10km plume assumption.

In terms of the comparison of the IASI SO<sub>2</sub> plume height against the S5P/TROPOMI layer height, for five coincident eruptions, we report that when comparing with IASI-A, the mean S5P SO<sub>2</sub> height is 11.5 $\pm$ 2.4 km and IASI-A 11.8 $\pm$ 2.1km, while for IASI-B the mean heights are 11.5 $\pm$  2.5km and 11.7 $\pm$ 2.0km, respectively. We may hence conclude the IASI-A and IASI-B SO<sub>2</sub> heights reside within 2.5km from the S5P SO<sub>2</sub> height estimates, well within the currently acceptable ranges for this product.

#### We hence recommend the public release of this product.

<sup>&</sup>lt;sup>4</sup> http://stcorp.github.io/harp/doc/html/ingestions/index.html



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## APPENDIX

Table 6. Daily mean SO<sub>2</sub> height [km] for all eruptive days shown in Figure 30, bottom panel.

Eruptive volcano	Date	S5P SO <sub>2</sub> height [km]	IASI-A SO2 height [km]	IASI-B SO2 height [km]	
	20200710		7.9±0.8	8.61±3.53	
	20200711	8.71±1.48	7.87±0.96	8.17±2.03	
	20200712	7.6±1.69	n/a	9.0±0.0	
	20200713	8.33±1.83	8.35±1.0	9.33±0.47	
	20200714	10.6±1.95	7.95±0.85	7.79±0.76	
	20200715	13.79±2.6	9.88±1.52	10.0±1.41	
	20200716	8.4±2.74	9.9±1.37	10.0±3.21	
Nishinoshima	20200717	8.94±1.32	8.5±0.5	8.0±0.0	
	20210410	14.14±3.48	14.27±1.77	14.24±1.76	
LaSoufriere	20210411	16.0±3.43	16.07±1.68	15.45±1.78	
Taal	20200113	10.69±3.55	14.57±2.39	15.76±2.16	
	20190623	8.83±2.73	11.01±2.61	10.7±2.37	
	20190624	9.65±1.92	10.9±1.94	10.89±2.27	
	20190625	9.23±2.07	10.88±1.74	10.84±1.72	
	20190626	9.2±2.6	11.5±2.02	11.31±1.77	
	20190627	10.05±2.92	11.95±2.47	11.96±2.54	
	20190628	9.56±2.99	11.95±2.46	12.04±2.59	
	20190629	10.09±2.86	12.25±2.67	12.08±2.61	
	20190630	9.18±2.91	12.09±2.78	12.14±2.84	
	20190701	8.06±3.01	12.26±2.19	12.27±2.25	
	20190702	8.94±3.84	11.93±1.67	12.16±2.03	
	20190703	9.42±3.38	n/a	11.97±1.95	
	20190704	10.03±3.4	n/a	11.98±2.08	
	20190705	12.63±3.92	n/a	11.95±2.02	
	20190706	14.38±5.14	n/a	11.95±2.01	
	20190707	15.37±4.9	n/a	11.9±2.07	
	20190708	13.76±2.73	12.12±2.12	12.05±2.12	
	20190709	13.76±2.4	12.32±2.27	12.31±2.31	
	20190710	13.81±2.38	12.44±2.24	12.35±2.28	
	20190711	12.71±1.76	12.47±2.19	12.35±2.27	
	20190712	14.23±2.69	12.52±2.33	12.52±2.42	
	20190713	14.52±2.42	12.94±2.5	12.85±2.53	
	20190714	12.07±2.46	13.19±2.61	13.14±2.6	
	20190715	11.77±1.06	13.2±2.81	13.17±2.84	
Raikoke	20190716	12.15±1.46	13.73±2.77	13.6±2.91	



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	20190717	10.0±1.14	14.18±3.0	13.72±2.94
	20190718	13.21±1.32	14.2±2.94	13.95±2.93
	20190719	13.31±2.08	14.44±2.87	14.12±2.92
	20190720	13.43±3.43	14.18±2.8	14.2±2.84
	20210221	13.76±2.47	11.1±1.66	10.63±1.87
	20210223	17.02±0.89	10.96±1.73	11.08±1.91
	20210304	13.67±3.43	11.16±2.66	11.43±2.72
Mt Etna	20210401	9.89±0.75	7.27±0.75	7.36±0.72